## "Tidal evolution of close-in planets using a Maxwell visco-elastic rheology"

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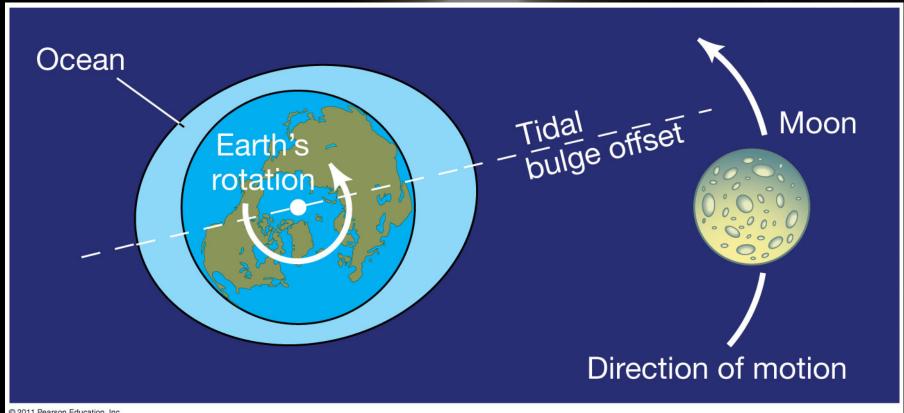
(Univ. Aveiro / IMCCE - Obs. Paris)

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Porto, Portugal

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## Tidal effects



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# Global Picture:

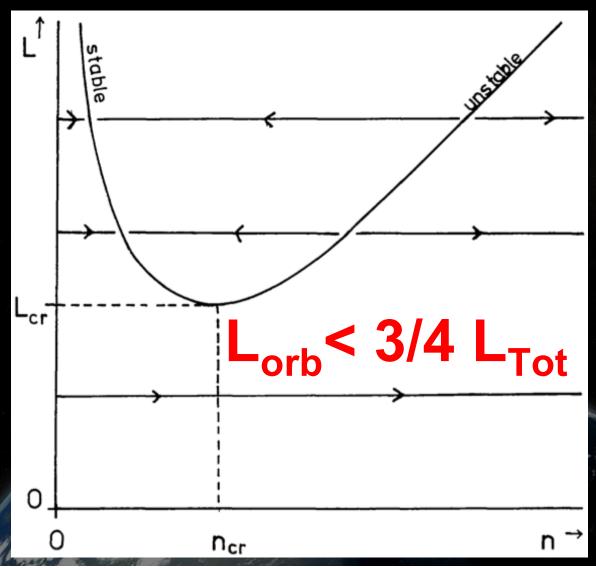
Poincaré (1898)

Landau & Lifshitz (1960)

Hut (1980)

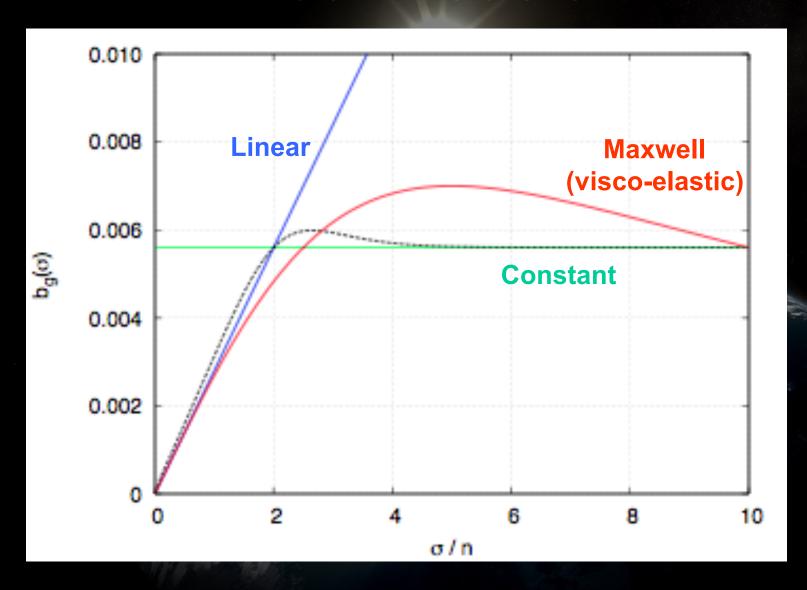
P<sub>orb</sub> = P<sub>rot</sub>
pependicular
axis (ε=0)

circular orbits (e=0)

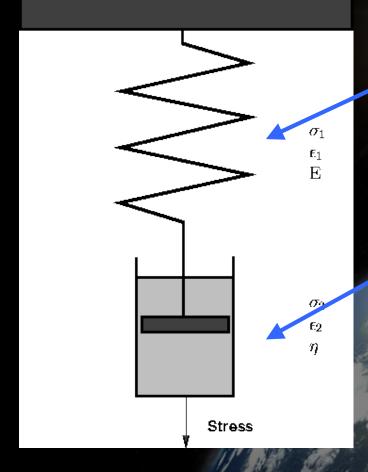


$$L_{cr} = 4 \left\{ \frac{1}{27} G \frac{M^3 m^3}{M+m} (I_1 + I_2) \right\}^{1/4}$$

## Tidal models



## Maxwell model (1867)



#### elastic

stress

$$\sigma = E \epsilon$$

strain

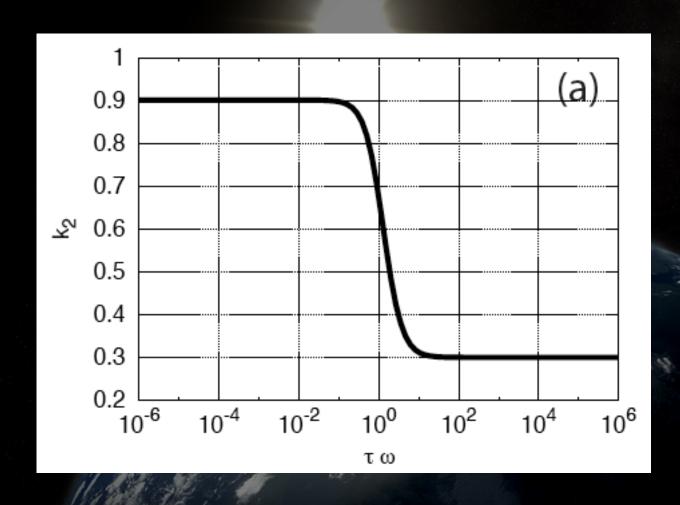
#### viscous

$$\sigma = \eta \, d\epsilon/dt$$

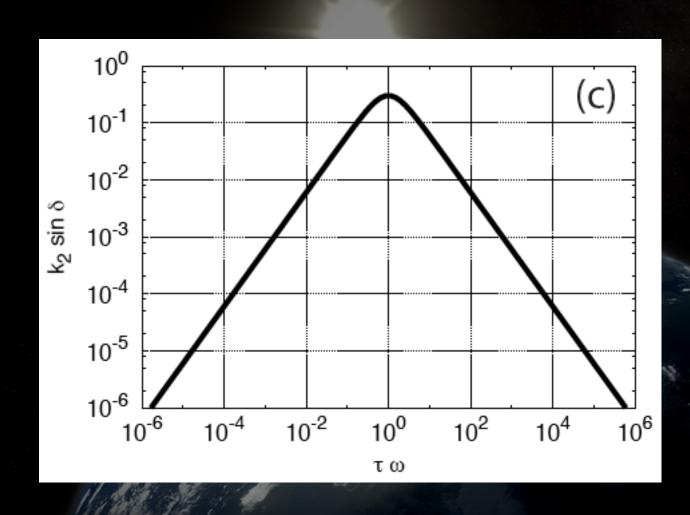
$$\frac{d\epsilon_{\text{Total}}}{dt} = \frac{d\epsilon_D}{dt} + \frac{d\epsilon_S}{dt} = \frac{\sigma}{\eta} + \frac{1}{E} \frac{d\sigma}{dt}$$

$$\frac{1}{E}\frac{d\sigma}{dt} + \frac{\sigma}{\eta} = \frac{d\epsilon}{dt}$$

## Deformation

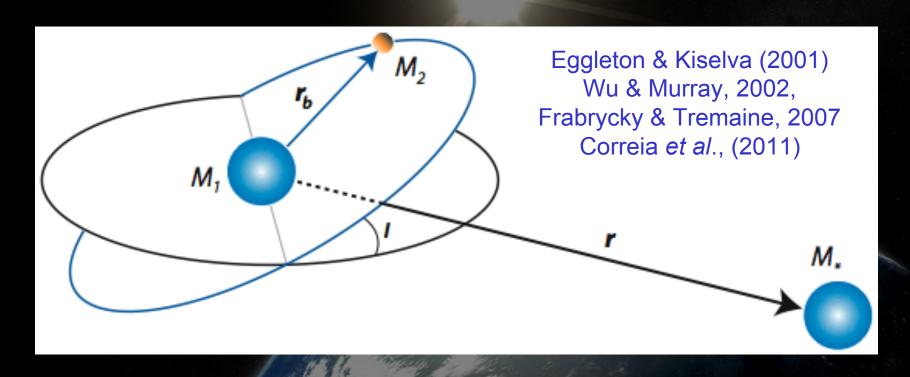


## Dissipation



#### HD 80606

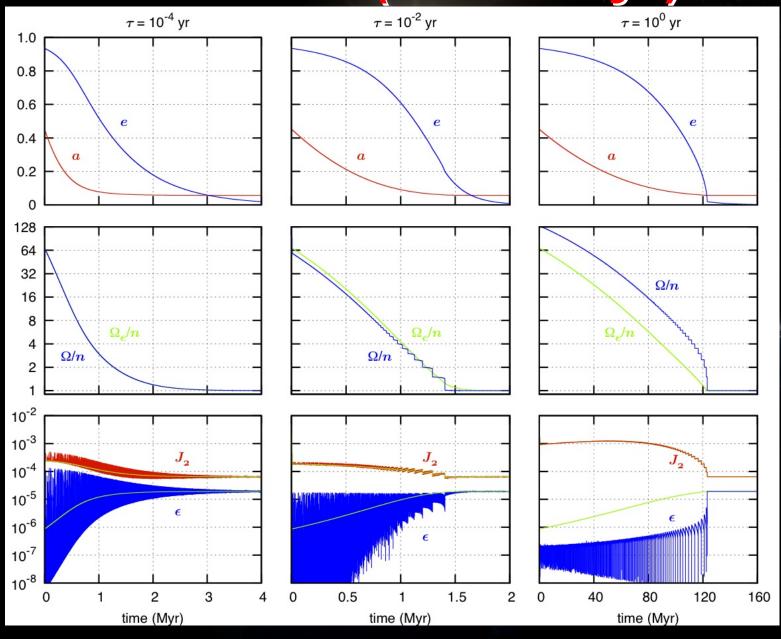
#### Naef et al, 2001 ( $a_p = 0.45 \text{ AU}, e_p = 0.92$ )



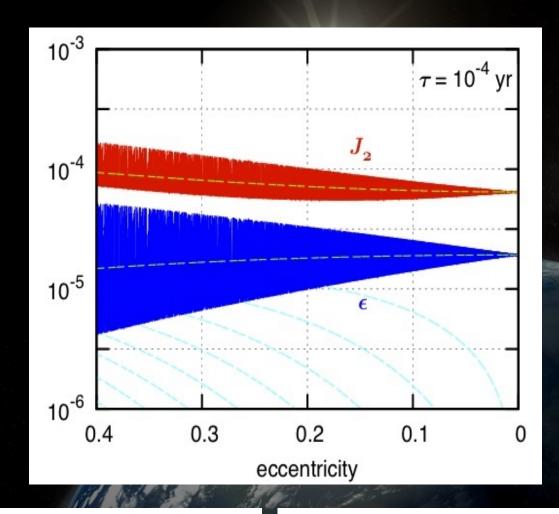
Kozai effect: (1 > 39°)

$$\sqrt{1 - e_1^2} \cos I = h_1 = Cte$$

## HD 80606 (n $\sim 10^{-2}$ yr)



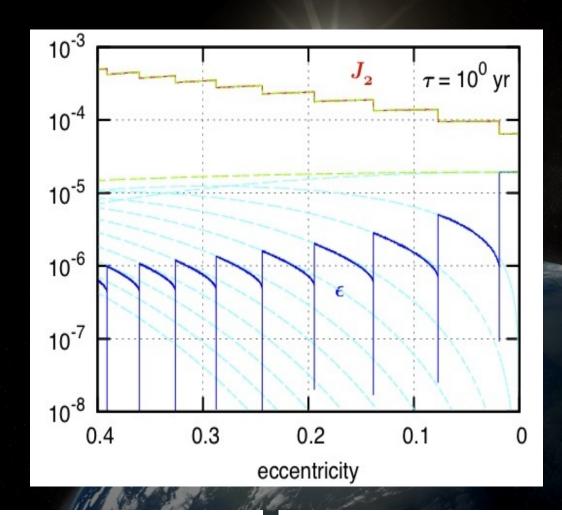
#### HD 80606 - deformation (t << 1/n)



$$\langle \epsilon \rangle_M = \frac{k_{\rm f}}{4} \frac{m_0}{m} \left(\frac{R}{a}\right)^3 (1 - e^2)^{-3/2}$$

$$\langle \epsilon \rangle_{M} = \frac{k_{\rm f}}{4} \frac{m_{0}}{m} \left(\frac{R}{a}\right)^{3} (1 - e^{2})^{-3/2} \left( \langle \epsilon_{p} \rangle_{M} \right) = \beta_{2p} = \frac{k_{\rm f}}{4} \frac{m_{0}}{m} \left(\frac{R}{a}\right)^{3} X_{2p}^{-3,2}(e)$$

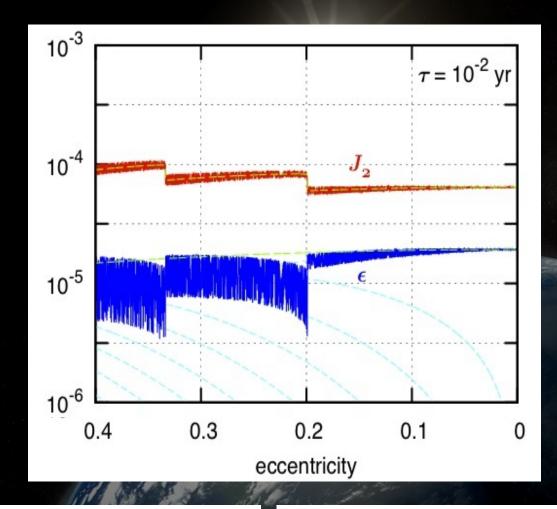
#### HD 80606 - deformation $(\tau >> 1/n)$



$$\langle \epsilon \rangle_M = \frac{k_{\rm f}}{4} \frac{m_0}{m} \left(\frac{R}{a}\right)^3 (1 - e^2)^{-3/2}$$

$$\langle \epsilon \rangle_M = \frac{k_{\rm f}}{4} \frac{m_0}{m} \left(\frac{R}{a}\right)^3 (1 - e^2)^{-3/2} \qquad \left\langle \epsilon_p \right\rangle_M = \beta_{2p} = \frac{k_{\rm f}}{4} \frac{m_0}{m} \left(\frac{R}{a}\right)^3 X_{2p}^{-3,2}(e)$$

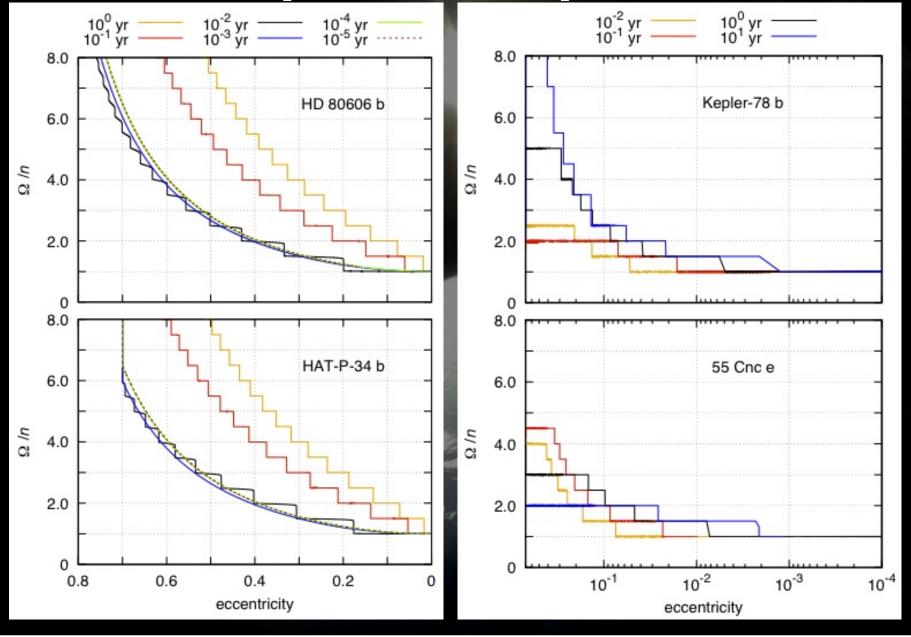
#### HD 80606 - deformation ( $\tau \sim 1/n$ )



$$\langle \epsilon \rangle_{M} = \frac{k_{\rm f}}{4} \frac{m_{0}}{m} \left(\frac{R}{a}\right)^{3} (1 - e^{2})^{-3/2} \left( \langle \epsilon_{p} \rangle_{M} \right) = \beta_{2p} = \frac{k_{\rm f}}{4} \frac{m_{0}}{m} \left(\frac{R}{a}\right)^{3} X_{2p}^{-3,2}(e)$$

$$\left\langle \epsilon_p \right\rangle_M = \beta_{2p} = \frac{k_{\rm f}}{4} \frac{m_0}{m} \left(\frac{R}{a}\right)^3 X_{2p}^{-3,2}(e)$$

## Hot-Jupiters / super-Earths



#### Conclusions

- We replaced a Fourier series by a rheological law of deformation of the potential. This allow us to simultaneously take into account deformation and dissipation.
- We no longer need to truncate the series for high eccentricities and we avoid a large number of terms. The deformation is also valid for all tidal regimes.
- Spin-orbit resonances arise naturally when the deformation time-scale is longer than the orbital period. The stability increases with the deformation time-scale.
- For rocky planets, the spin-orbit resonances are possible for very low values of the eccentricity, so we expect that some of these planets are not synchronous with the star.